# Turbulent transport in stably stratified atsmospheres

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July 9, 2010

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## Turbulent transport in stably stratified atmospheres

Transport in the ocean, in planetary atmospheres

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 Transport of chemical elements/angular momentum in stellar radiative zones



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- Some remarks on the mixing length theory
- The G.I. Taylor diffusion theory
- Vertical transport of chemical elements in homogeneous stably stratified turbulence
- Vertical transport of chemical elements in stellar radiative zones

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### Based on an analogy with molecular diffusion in perfect gas

- Small departure from thermodynamic equilibrium  $\Rightarrow \nu = \frac{1}{3}\lambda v_{th}$
- Prandtl's assumption  $\nu_T = \ell_m v = \ell_m^2 |\frac{dU}{dz}|$

#### But ...

- while λ ≪ L, there is no such scale separation in general in turbulence : for example, in free shear flows, ℓ ~ L.
- the theory is incomplete :  $\nu_T = \ell_m v$  but  $\ell_m$  and v are unknown.

## Does it work anyway ?

Yes, for one-dimensional free shear flows U(z)

$$\ell_m = cL$$

where 0.07  $\leq c \leq$  0.18 depending on the type of shear flows

For this one length scale L and one time scale  $1/\frac{dU}{dz}$  problem, this success can be attributed to dimensional analysis :

$$\ell \sim L \qquad \sqrt{\langle u^2 \rangle} / \ell \sim \frac{dU}{dz}$$
$$\langle uw \rangle | \sim \sqrt{\langle u^2 \rangle} \sqrt{\langle w^2 \rangle} \sim \langle u^2 \rangle \sim |L^2 \left(\frac{dU}{dz}\right)^2 | \sim \nu_T |\frac{dU}{dz}$$

#### But ... if the fluid is stably stratified

the buoyancy force introduces another time scale 1/N

$$\Rightarrow c = f(N/\frac{dU}{dz})$$
 with  $f$  unknown

## G.I. Taylor's turbulent diffusion model

### Vertical displacement of single particule

$$\frac{dz}{dt} = W(t) \quad \Rightarrow \quad z - z_0 = \int_0^t W(s) ds$$

$$(z-z_0)W(t) = W(t)\int_0^t W(s)ds$$
$$\frac{1}{2}\frac{d(z-z_0)^2}{dt} = \int_0^t W(t)W(s)ds$$

### Ensemble average over many particules < .. >

$$rac{1}{2}rac{d < (z-z_0)^2 >}{dt} = \int_0^t < W(t)W(s) > ds$$

The mean displacement depends on the Lagrangian autocorrelation :

$$R_{t,s} \;\; = \;\; rac{< W(t) W(s) >}{< W(t)^2 >^{1/2} < W(s)^2 >^{1/2}}$$

## G.I. Taylor's turbulent diffusion II

### Statistical stationarity

$$\frac{\langle W(t)W(s) \rangle}{\langle W(t)^{2} \rangle^{\frac{1}{2}} \langle W(s)^{2} \rangle^{\frac{1}{2}}} = \frac{\langle W(t)W(t+s-t) \rangle}{\langle W^{2} \rangle} = R_{s-t}$$

$$\frac{1}{2}\frac{d \langle (z-z_{0})^{2} \rangle}{dt} = \langle W^{2} \rangle \int_{0}^{t} R_{s-t}ds = \langle W^{2} \rangle \int_{0}^{t} R_{u}du$$

$$\langle (z-z_{0})^{2} \rangle (T) = 2 \langle W^{2} \rangle \int_{0}^{T} \left( \int_{0}^{t} R_{u}du \right) dt$$

### The Lagrangian correlation time $\tau_L$

As 
$$\tau_L = \lim_{t \to \infty} \int_0^t R_u du$$
 is finite  $\Rightarrow$   
 $< (z - z_0)^2 > (T) = 2 < W^2 > \tau_L T$  for  $T \gg \tau_L$   
Brownian like diffusion with  $D = < W^2 > \tau_L$ 

## Vertical transport of chemical elements in homogeneous stably stratified turbulence Models

The two successive transport regimes predicted by the Langevin-type model of Pearson et al. (1983)

$$<\delta z^2>=rac{}{N^2}+crac{}{N}t$$

- ► Adiabatic regime : suppression of the vertical dispersion by the buoyancy force < δz<sup>2</sup> >= ≤w<sup>2</sup>>/M<sup>2</sup>
- ► Diabatic regime : vertical diffusion induced by thermal diffusivity  $<\delta z^2>=c\frac{<w^2>}{N}t$

According to Lindborg & Fedina (2008),  $D = \epsilon_P N^2$ , where  $\epsilon_P$  is the rate of potential energy dissipation through molecular diffusion.

## Vertical transport of chemical elements in homogeneous stably stratified turbulence Numerical simulations



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# Vertical transport of chemical elements in stellar radiative zones

#### Observational constraints

- Abundance anomalies in the sun, Am-Fm stars can be accounted for by very small effective diffusivities (Michaud & Zahn 1998) e.g. D = 2000cm<sup>2</sup>/s in the solar radiative zone
  - $\Rightarrow$  Small length scales and/or small velocity scales

#### The dynamical role of thermal diffusion

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- reduces the amplitude of the buoyancy force
- dissipate kinetic energy
- changes the equilibrium level of fluid elements

## Two different regimes depending on the amplitude of vertical velocities (Lignières et al. 2005)



 $D = \ell_{B\kappa} w < \kappa$  as expected from the observations